

Observatoire Côte d'Azur - 18th March 2014

Comparing atmospheric chemical disequilibrium of Earth and Mars to detect the traces of Life

Eugenio Simoncini, John R. Brucato

Astrophysical Observatory of Arcetri - INAF, Firenze, Italy

Tommaso Grassi

Starplan, Copenhagen, Denmark / University of Rome - La Sapienza, Roma, Italy



Thermodynamics,
Disequilibrium and
Evolution (TDE) – NAI
Focus Group



Who I am



- * *Co-Chair of the “Thermodynamics, Disequilibrium and Evolution” NAI Focus Group*
- * Astrophysical Observatory of Arcetri, INAF, Firenze, Italy
=> Atmospheric habitability; hydrothermal vents thermodynamics; life/habitability detection devices for space missions (May 2013)
- * Centre of Astrobiology, INTA-CSIC, Madrid, Spain (2012)
- * Max Planck for Biogeochemistry, Jena, Germany (2010 - 2012)

2010 PhD in Chemical Sciences (*Physical Chemistry; chemical oscillators; thermodynamics; excitable media*).

2006 Master Degree in Chemistry for Sustainable Development (*Environmental Chemistry; “evolutive” thermodynamics; environmental accounting for land, productions and bioarchitecture*).

2004 Bachelor Degree in Chemistry (*Green Chemistry; Supercritical fluids; Ionic Liquid; efficiency management*).

(University of Siena, Italy)



Current projects

- Atmospheric chemical disequilibrium
- Life emergence by *chemiosmosis* in hydrothermal vents
- Mars surface and sub-surface disequilibrium
- Adsorption of molecules on minerals
- Planetary protection

Overview of this talk

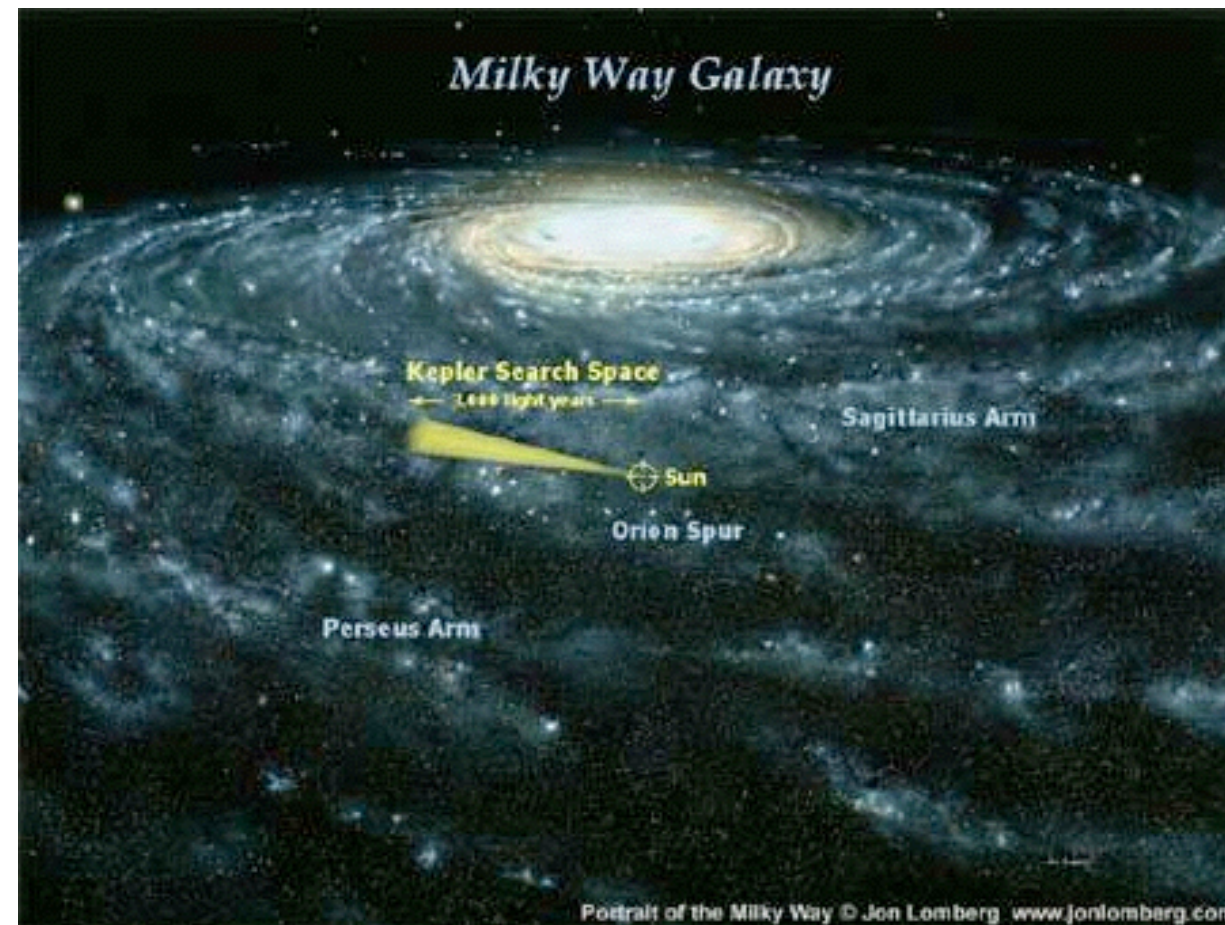
- What is chemical disequilibrium
- How to calculate disequilibrium in chemical processes
- Atmospheric Disequilibrium of Earth and Mars

Introduction

What is chemical disequilibrium, and why should we use it

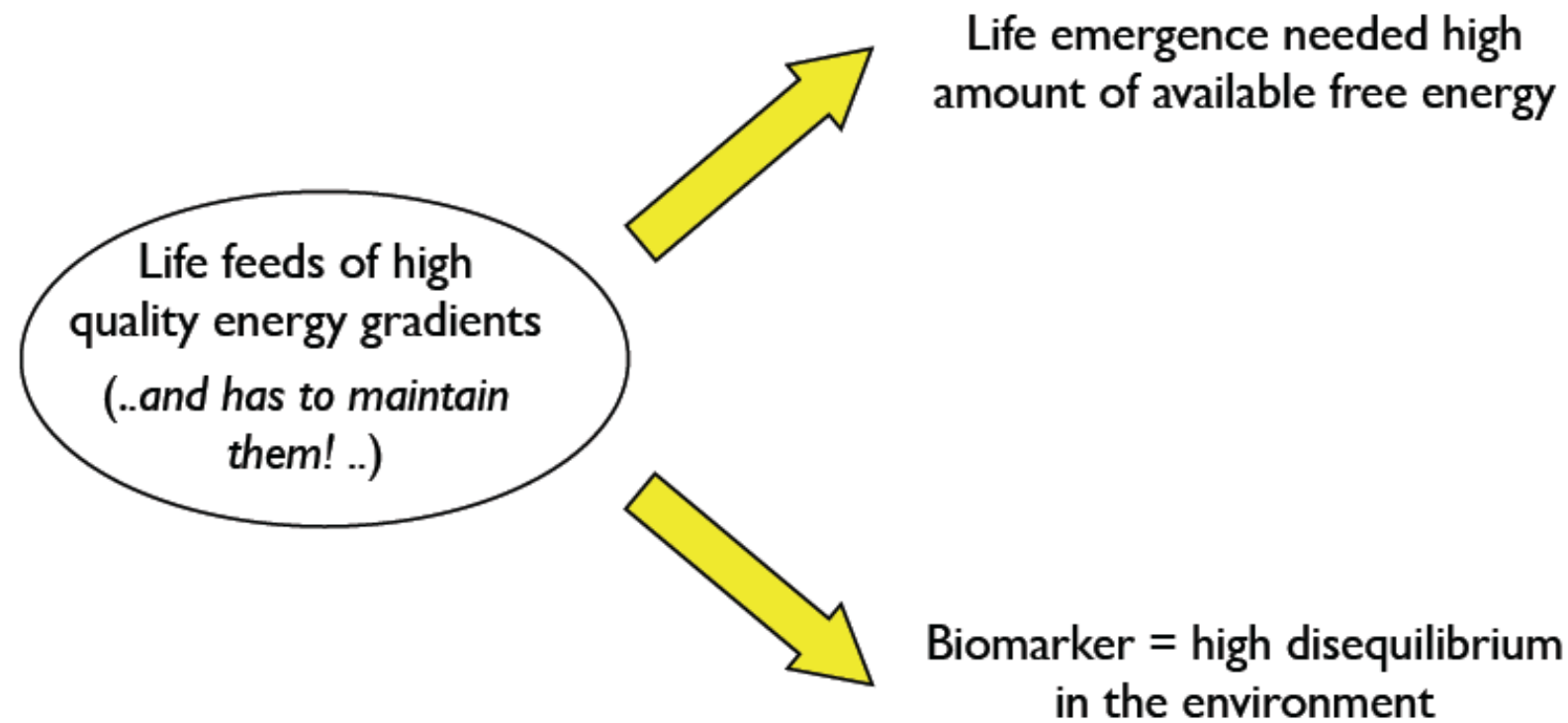
Life and Biomarkers

- no signs for complex molecules
- simple molecules spectra



Jim E. Lovelock, *Nature* 207, 568 (1965), about life:
“search for the presence of compounds in the planet’s atmosphere
that are incompatible on a long-term basis”

Life & Disequilibrium

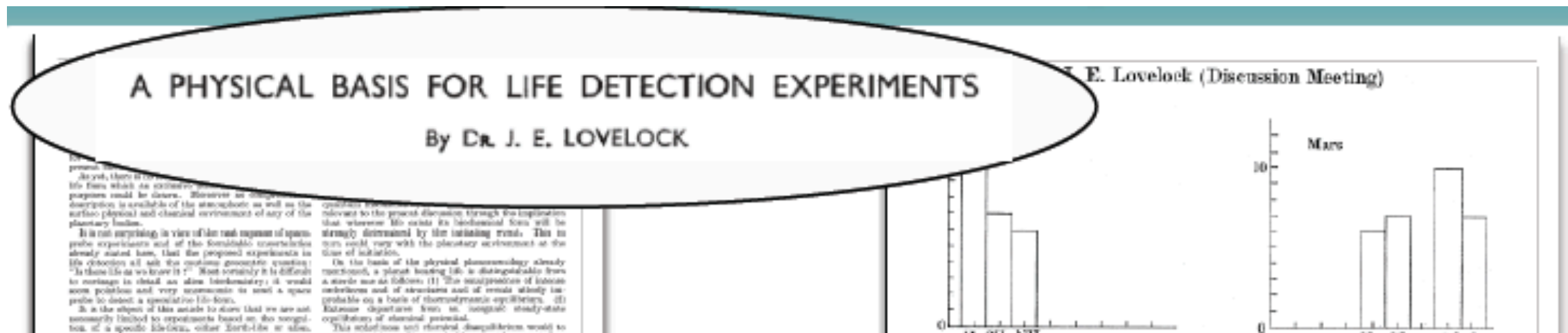


“The general struggle for existence of animate being is struggle for entropy, which becomes available through the transition of energy from the hot sun to the cold earth” (Boltzmann, 1886)

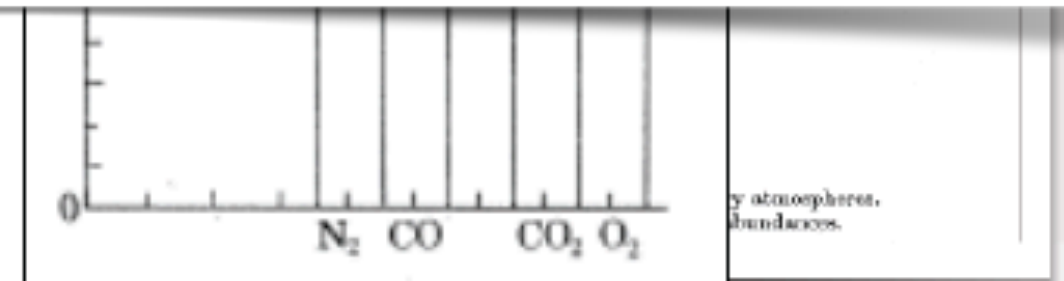
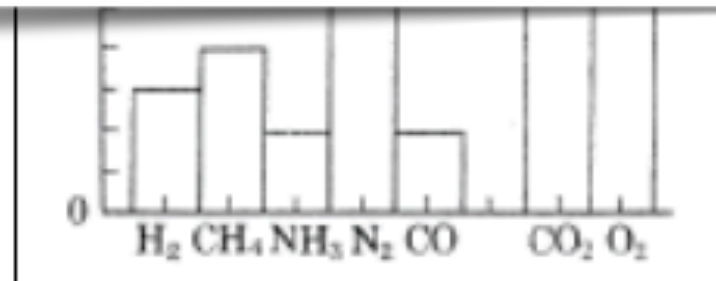
“Life feeds of high quality energy gradient” (Schrödinger, 1944)

“Once candidate disequilibria are identified, alternative explanations must be eliminated. Life is the hypothesis of last resort” (Sagan et al., 1993)

Atmospheric Chemical Disequilibrium

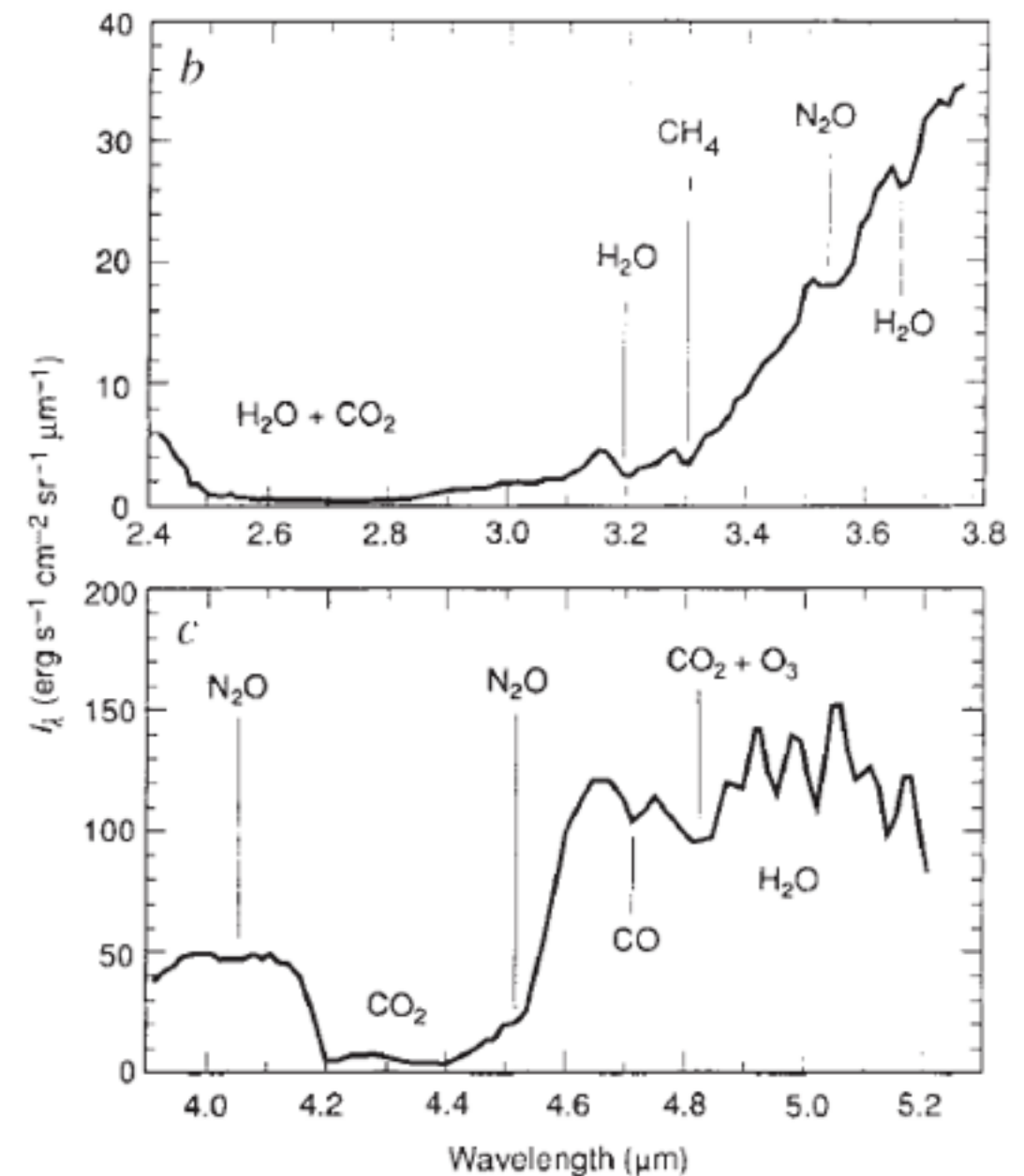
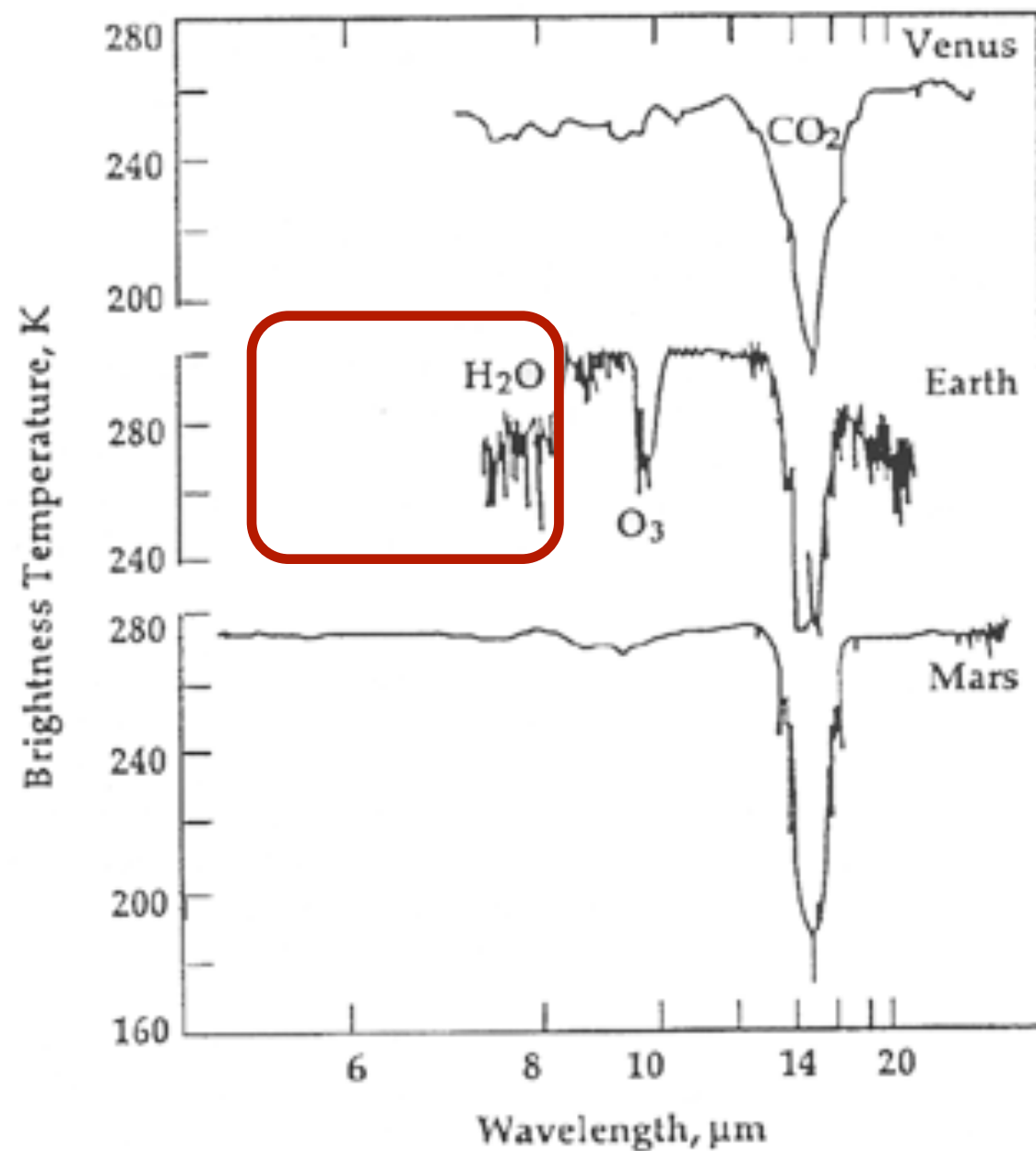


- * Quantify the extent of chemical disequilibrium
- * Consider other kind of chemical disequilibrium and different sources:
 - Fast vertical mixing vs. chemical kinetics (Hot Jupiters)
 - O_2/O_3 in Mars, Venus



Atmospheric Chemical Disequilibrium and Life

Earth



How to calculate (and compare) disequilibrium in chemical processes

The extent of chemical disequilibrium

In order to measure the extent of disequilibrium, we have to deal with the thermodynamics of non-equilibrium (irreversible) processes.

The distance of a system from its equilibrium condition (i.e. the measure of its irreversibility) is given by the entropy production within a system:

$$d_i S/dt$$

The extent of chemical disequilibrium



$$\frac{d_i S}{dt} = J \cdot X = \frac{d\xi}{dt} \cdot \frac{\alpha}{T}$$

Extent of reaction:

$$\xi(t) = \frac{[A]_0 - [A](t)}{\nu_A}$$

Chemical Affinity

$$\alpha(t) = - \left(\frac{\partial \Delta_r G(t)}{\partial \xi} \right)_{T,p}$$

It can be also written as:

$$\frac{d_i S}{dt} = R \cdot (R_f - R_r) \cdot \ln \left(\frac{R_f}{R_r} \right)$$

R_f = forward rate

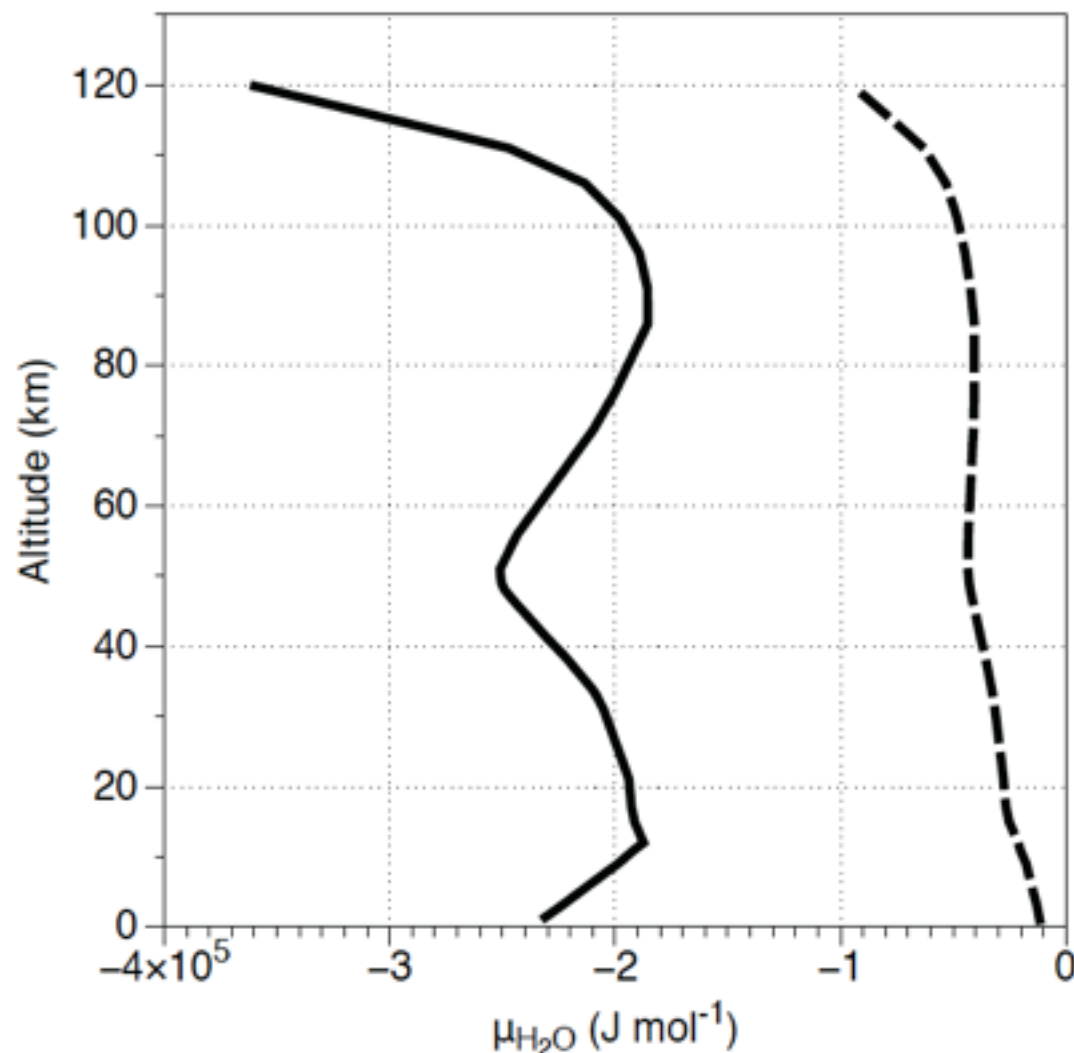
R_r = backward rate

Application to atmospheric disequilibrium

[The chemical potential for gas mixtures]



$$\mu_i(T, P, \chi_i) = \frac{T}{T_0} \mu_i^0 + \int_{p_0}^p V_{m,i}(T, p') dp' + T \int_{T_0}^T \frac{-h_i(T', p_0)}{T'^2} dT'$$



Key point:

$$\mu_i = \mu_i(t)$$

Simoncini E., Delgado-Bonal A., Submitted to Chem. Phys. Lett.

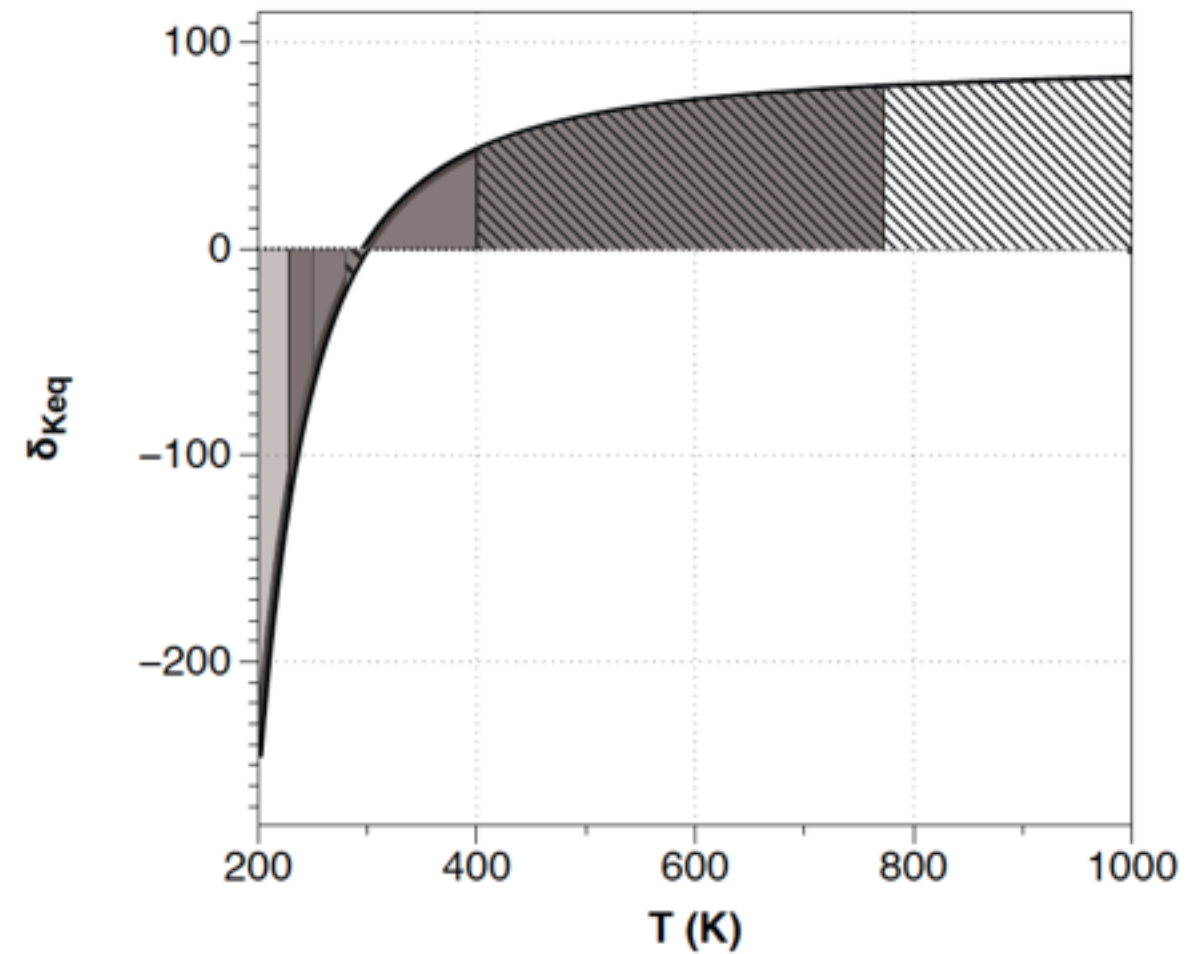
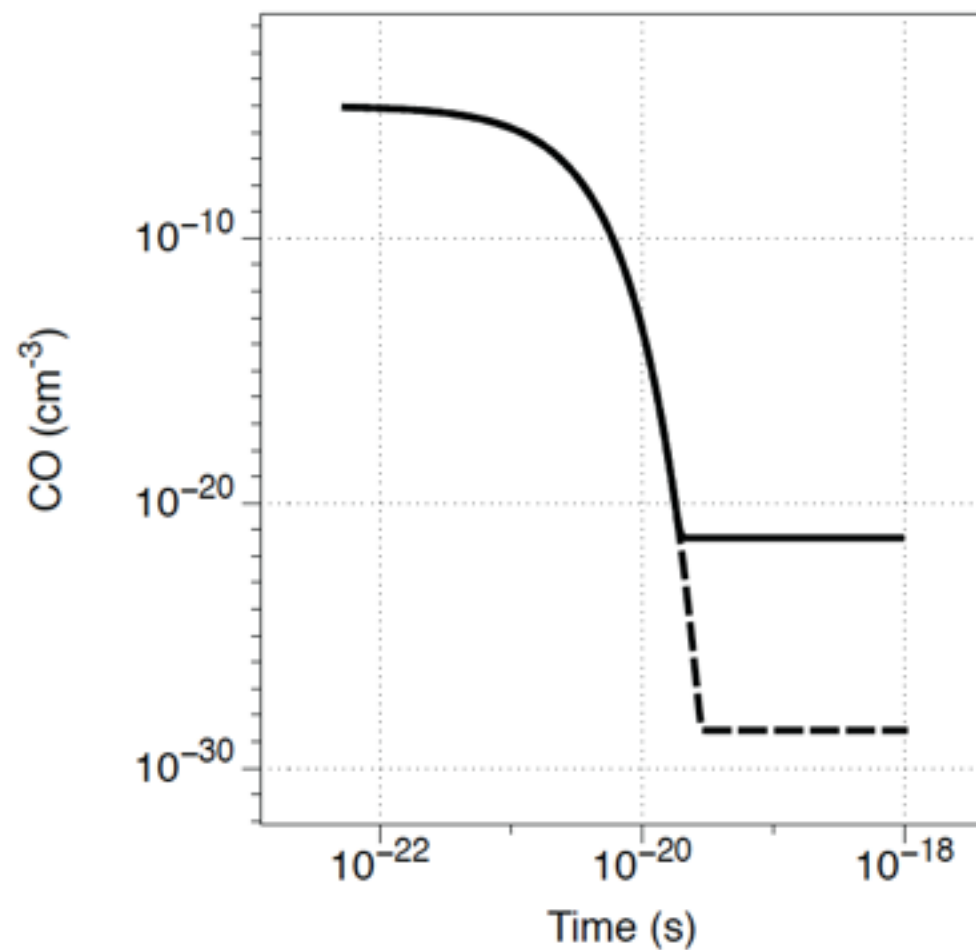
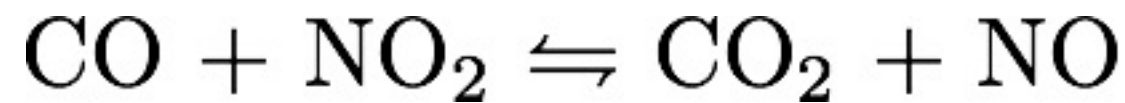
Kondepudi D., Prigogine I., Modern thermodynamics, 1998.

Chase, M. W. 1998, J. Phys. Chem. Ref. Data, Vol. 28.

NASA polynomials: http://www.me.berkeley.edu/gri_mech/data/nasa_plnm.html

[The chemical potential and the reaction kinetics]

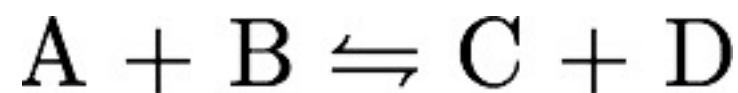
$$k_r = k_f \cdot \exp \left[\frac{\Delta_R G^0}{R \cdot T_0} + K_R(T) \right]$$



Atmospheric modeling with



$$\frac{d_i S}{dt} = R \cdot (R_f - R_r) \cdot \ln \left(\frac{R_f}{R_r} \right)$$



$$\frac{d_i S}{dt} = R(k_f[A]_t[B]_t - k_r[C]_t[D]_t) \ln \left(\frac{k_f[A]_t[B]_t}{k_r[C]_t[D]_t} \right)$$

→ A package to solve ODEs for the kinetics of hundreds of reactions
(like in atmospheres)

+ Calculate the entropy production

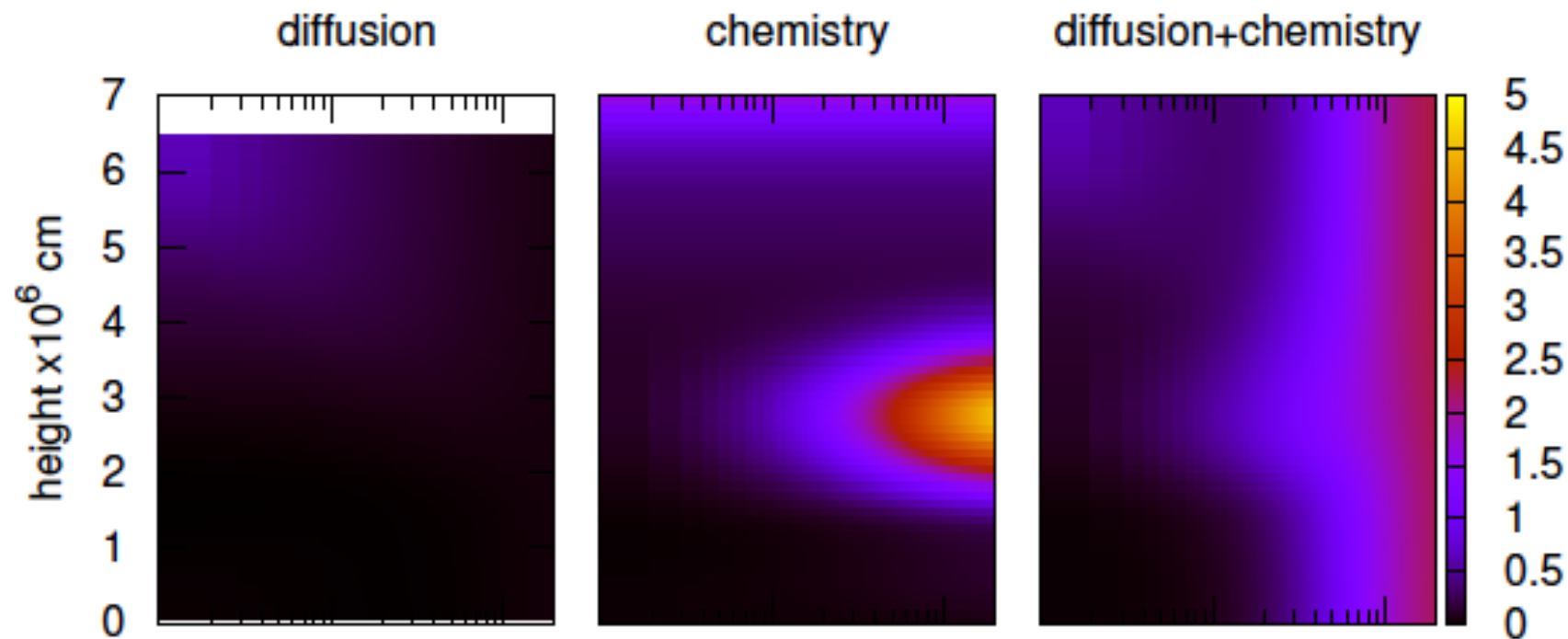


- Python pre-processor provides Fortran routines
- Creates modules from chemical network
- Dust evolution, cooling heating photoionization
- Large test suite
- Highly optimized, fast solvers
- Open source, bitbucket community
- Grassi T. et al., MNRAS 2014. doi:10.1093/mnras/stu114

www.kromepackage.org

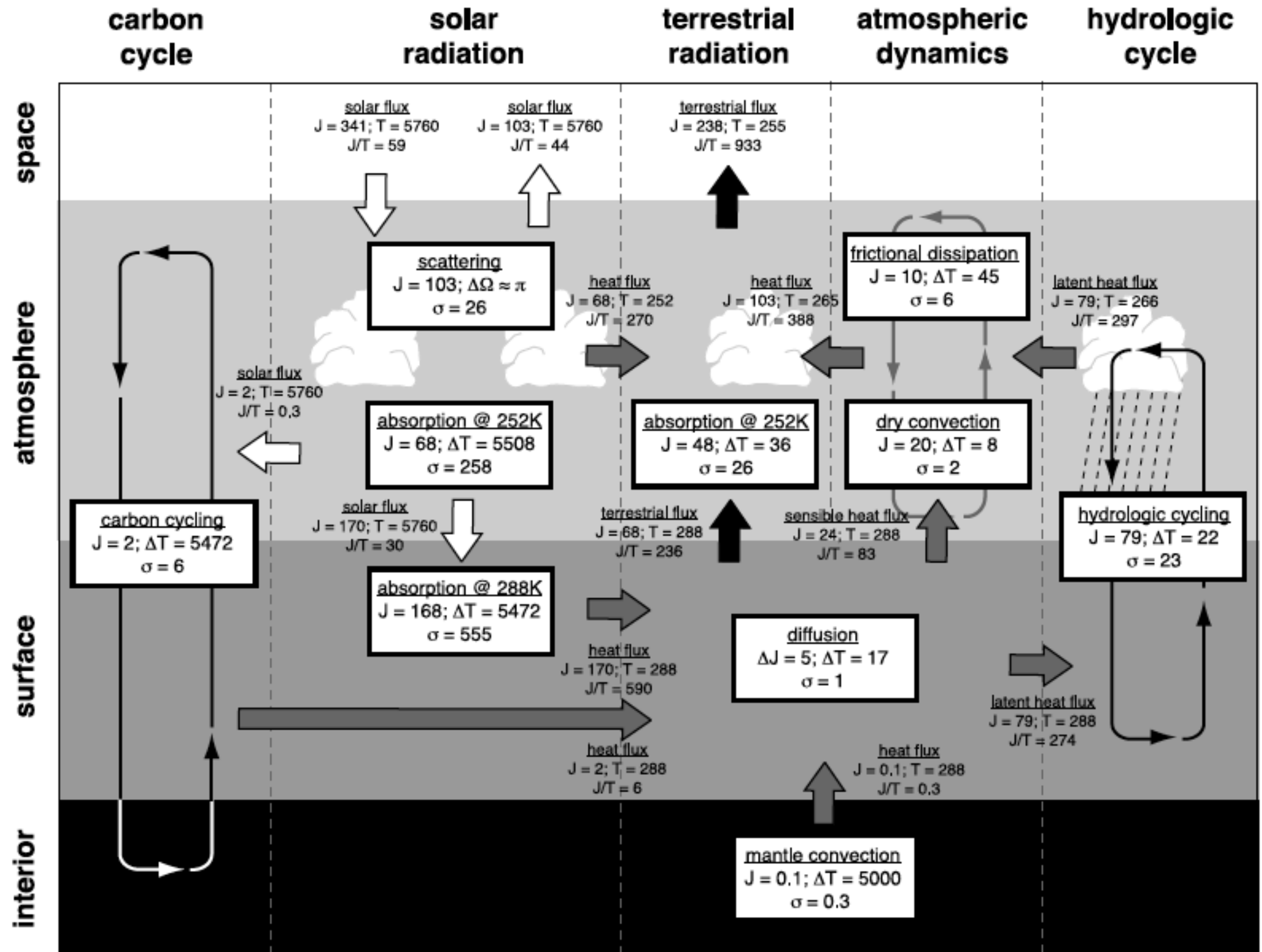
Diffusion and Kasting+80 chemical network (1D)

$$\frac{\partial n_{ij}}{\partial t} = \zeta_j \frac{\partial^2 n_{ij}}{\partial z^2}$$

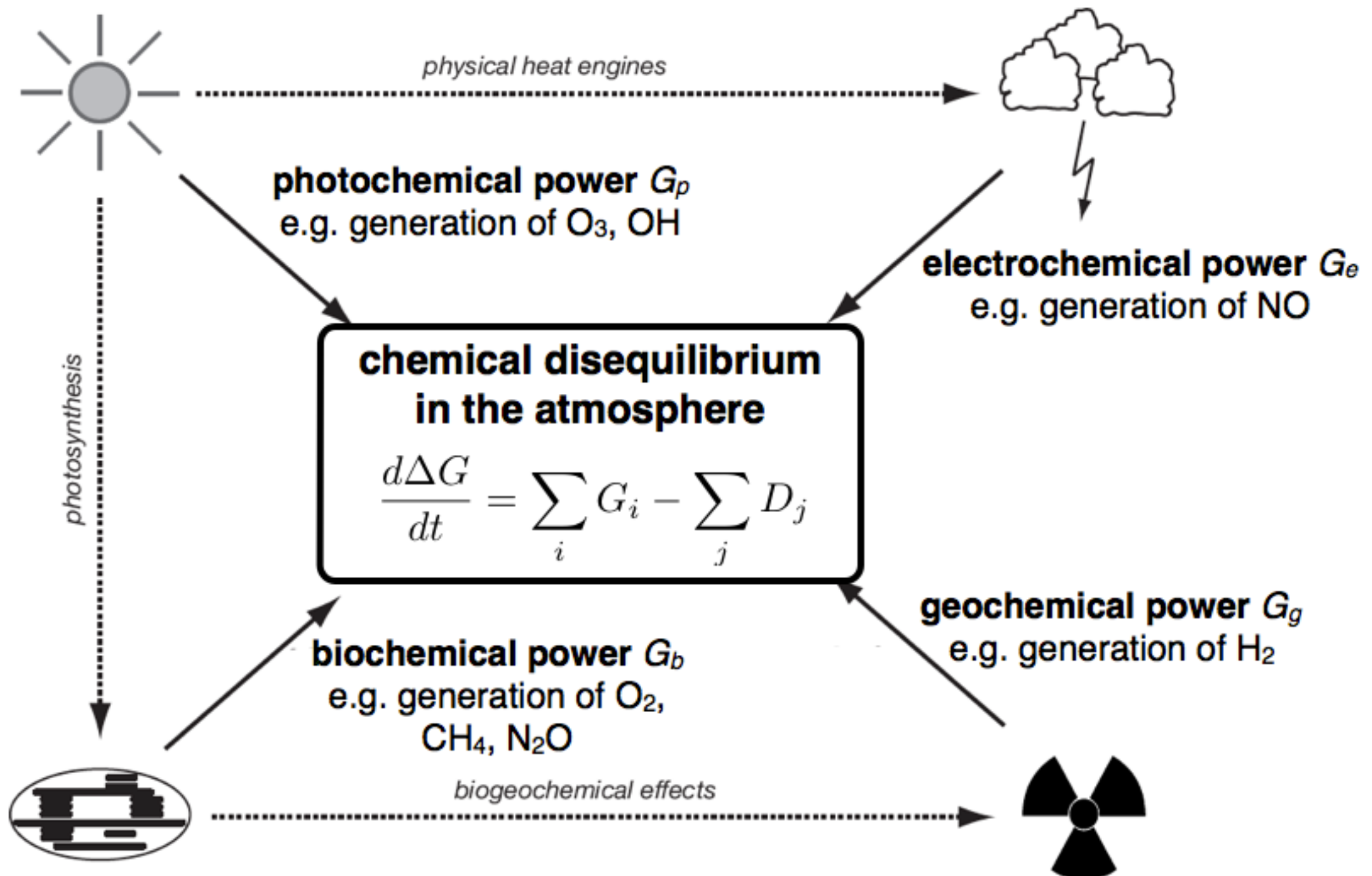


Earth Powers

$$J \sim W m^{-2}$$



Earth far from equilibrium



Earth Atmospheric Chemical Disequilibrium

Our first calculation:

- * Model: Kasting, J. F., and Donahue, T. M., J. Geophys. Res., 85,3255-3263. 1980;
- * 64 layers (~1km each);
- * Eddy diffusion;
- * Entropy production and the power dissipation:

$$\sigma = \frac{d_i S}{dt} \quad \frac{\sigma \times T}{A_{Earth}} \sim W m^{-2}$$

Simoncini & Grassi, submitted to OLEB.

Simoncini, Brucato, Grassi, in preparation

S. O. Danielache, E. Simoncini, Y. Ueno, Archean Atmospheres Modeled with the KROME Chemistry Package, JPGU 2014

Simoncini E., Virgo N., Kleidon A., Quantifying drivers of chemical disequilibrium: theory and application to methane in the Earth's atmosphere. Earth System Dynamics 4, 1-15, 2013.

Angerhausen D., Sapers H., Simoncini E., and coworkers, An astrobiological experiment to explore the habitability of tidally locked M-Dwarf planets, IAU 2013 Proceedings.

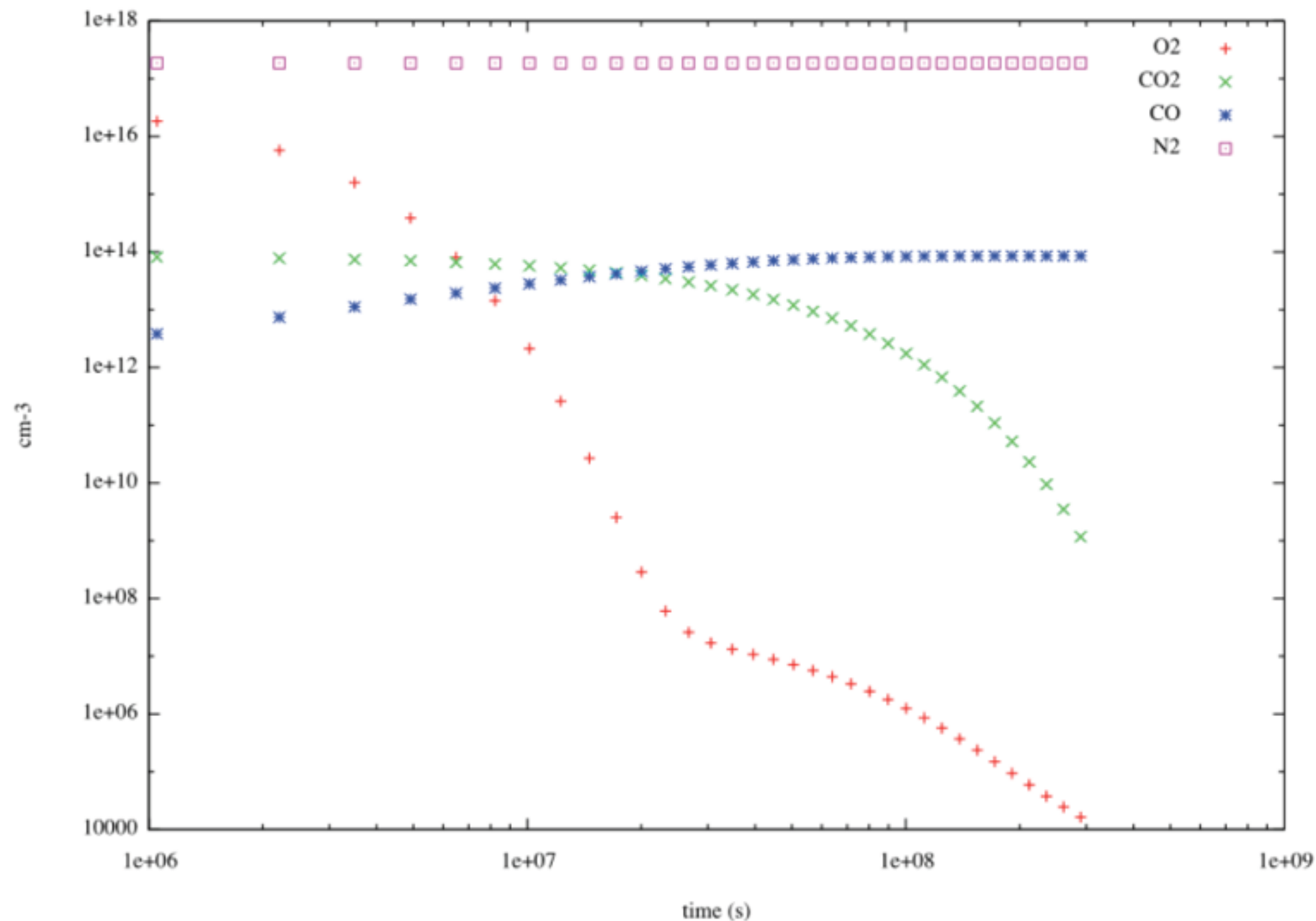
Earth Atmospheric Chemical Disequilibrium

Different runs of K-80 model:

- With and without photochemistry
- With and without eddy diffusion
- Low O₂, O₃, high CO₂ = pseudo pre-photosynthesis Earth

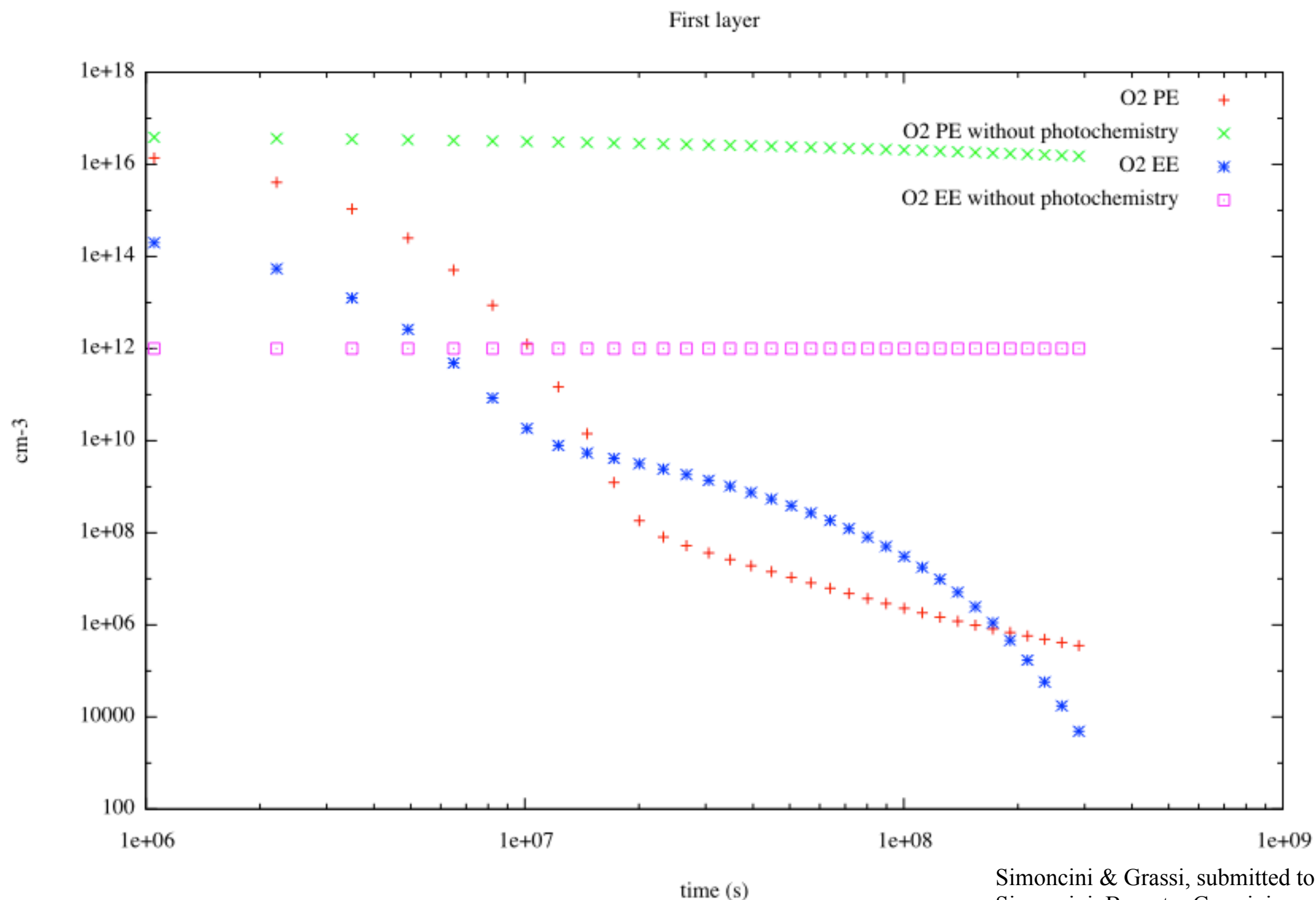
Earth Atmospheric Chemical Disequilibrium

Column model + KROME: species evolution

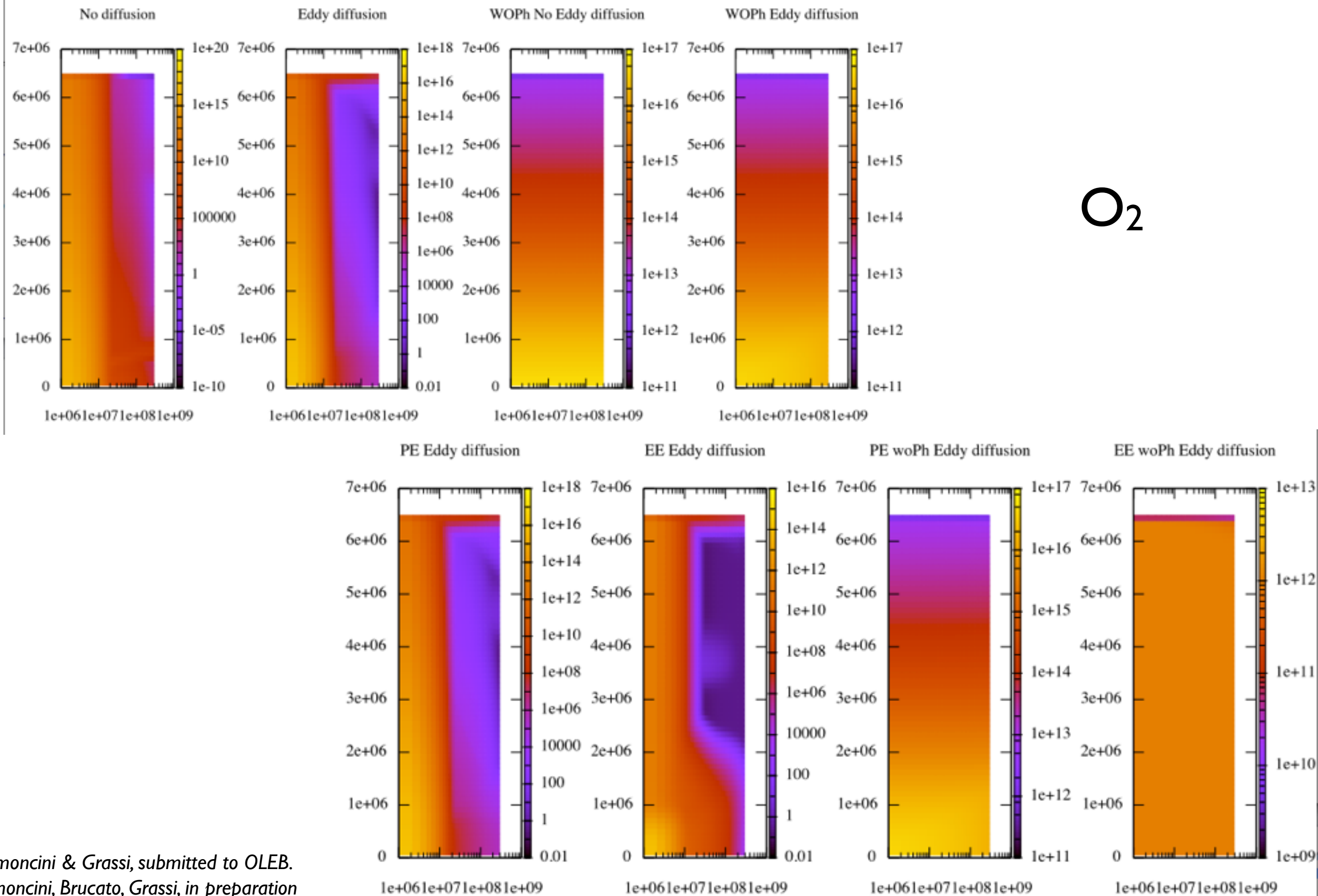


Earth Atmospheric Chemical Disequilibrium

Column model + KROME: species evolution



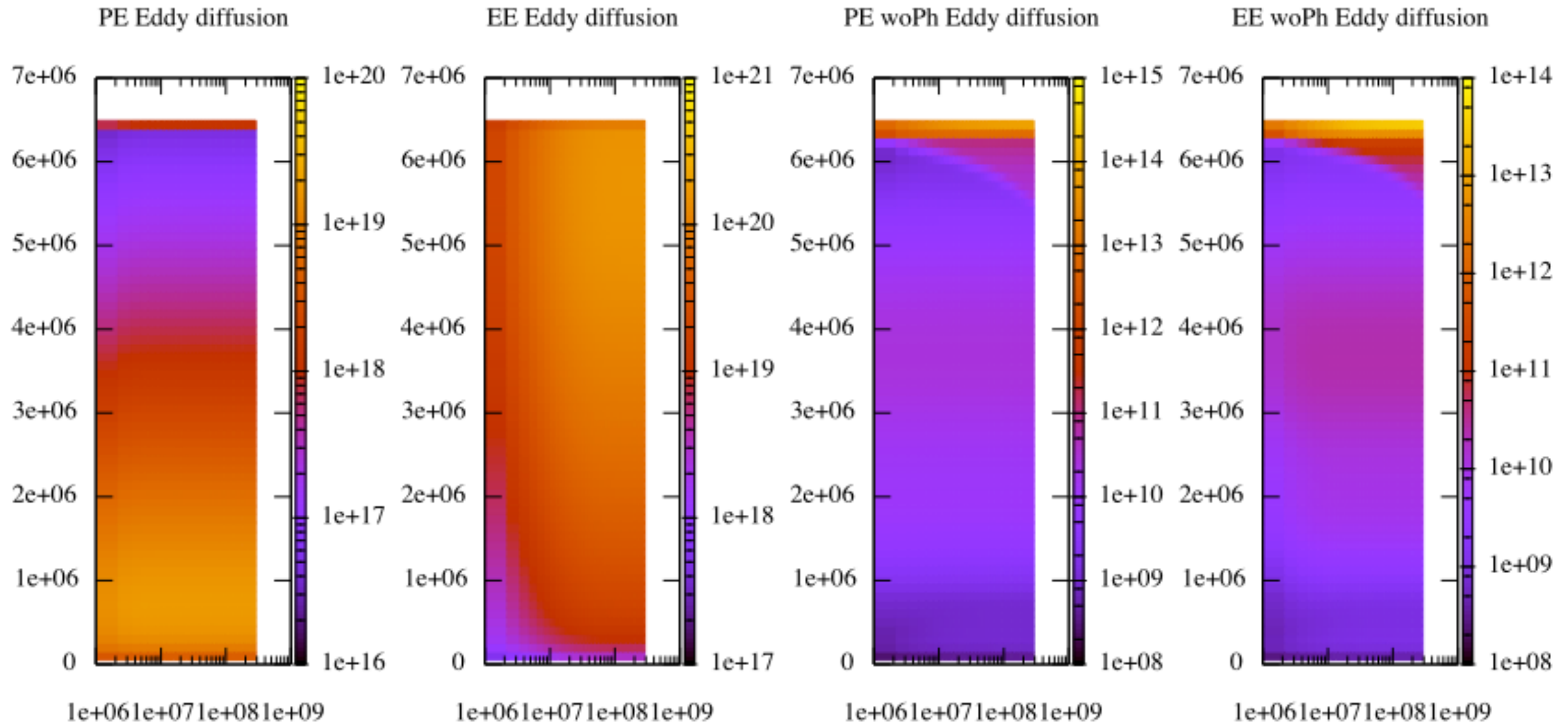
Earth Atmospheric Chemical Disequilibrium



O_2

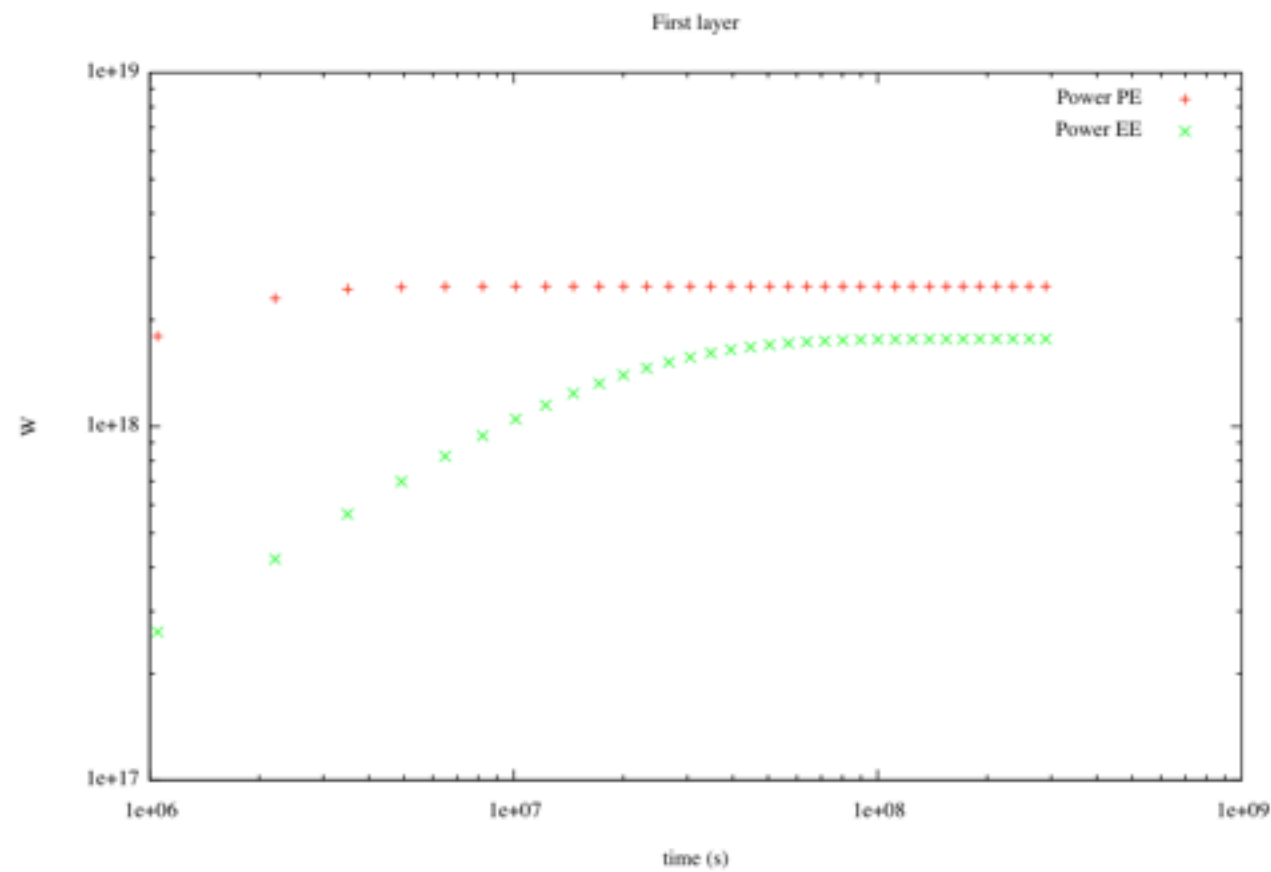
Earth Atmospheric Chemical Disequilibrium

The structure changes!

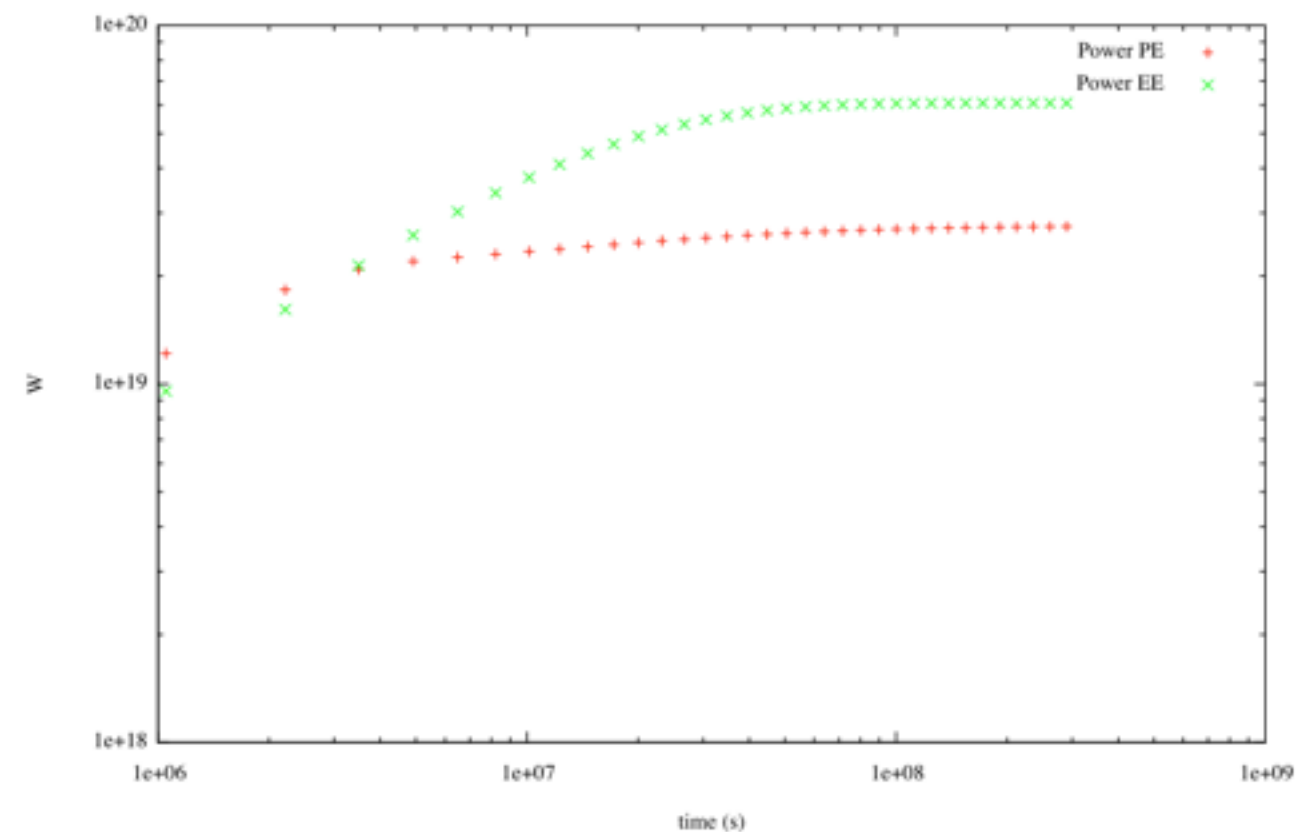


time (s)

Earth Atmospheric Chemical Disequilibrium



Layer 1



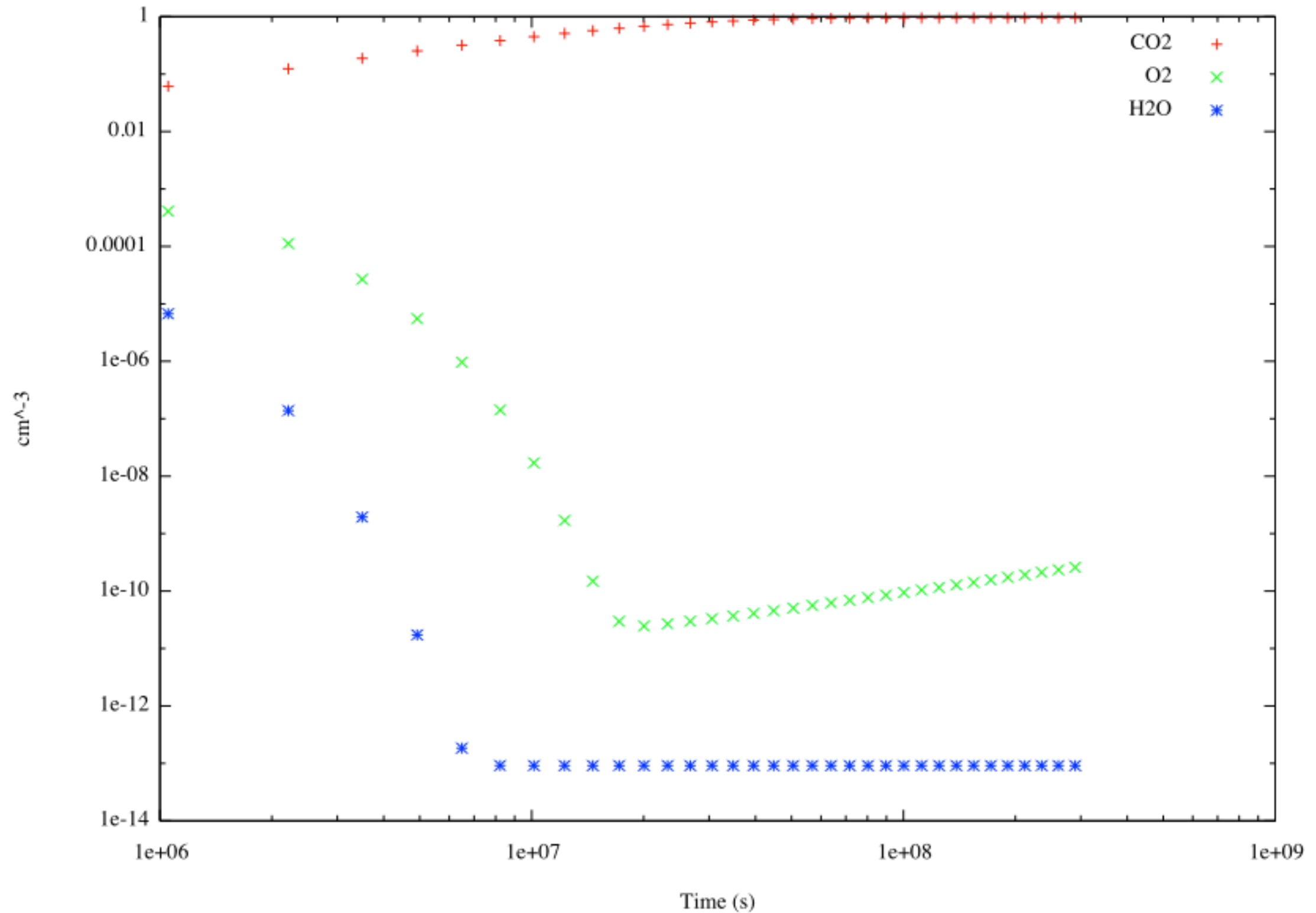
Layer 64

Mars Atmospheric Chemical Disequilibrium

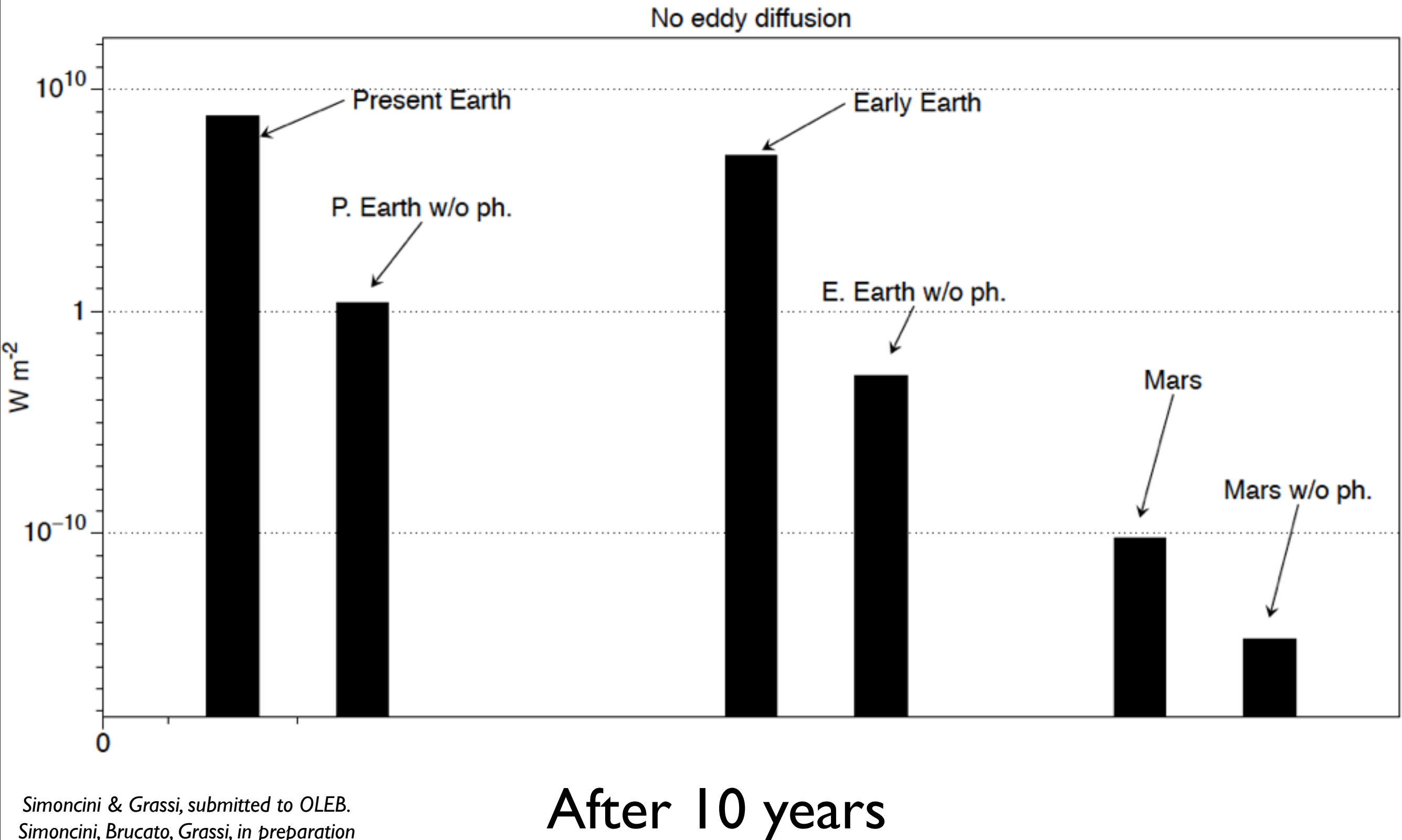
Our first calculation:

- * Model: Zahnle K., Haberle R. M., Catling D. C., Kasting J. F., JGR 113, 2008
- * 142 layers;
- * No diffusion
- * With and without photochemistry
- * Entropy production and the power dissipation.

Mars Atmospheric Chemical Disequilibrium



Atmospheric Chemical Disequilibrium: Mars vs Earth



Atmospheric Chemical Disequilibrium: Mars vs Earth

PE: 3.34 E23 W

PE no photochemistry: 1.56 E3 W

EE: 2.68 E6 W

EE no photochemistry: 1.17 E2 W

Mars: 8.70 E3 W

Mars no photochemistry: 2.59 E-1 W

Atmospheric Chemical Disequilibrium: Mars vs Earth

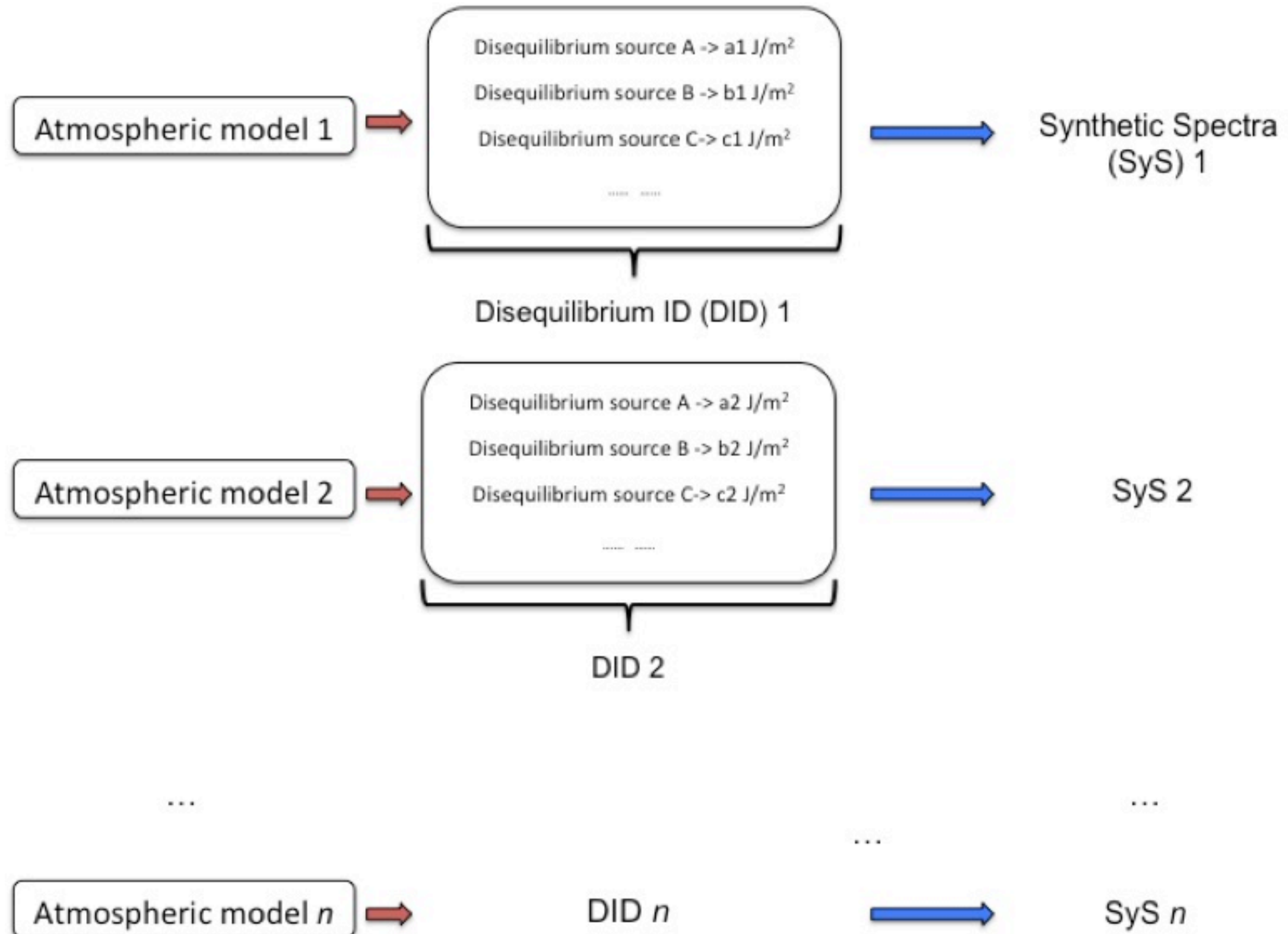
Results:

- * The effect on the atmospheric chemical disequilibrium attributable to the biosphere is very higher than the one due to photochemistry (some orders of magnitude).
- * Life and photochemistry are strongly linked in their effects. However, we found that the effect of photochemistry is enhanced by the biosphere.
- * Mars disequilibrium is more than 10 orders of magnitude lower compared to present and pre-photosynthesis Earth!

Perspectives

- Comparison of entropy production between planets (better link with the concept of “habitability”)
- Comparison between different power sources (photochemistry, fast vertical mixing, geological degassing, and life as the last option). Give a range of the maximum energy that can be dissipated by each process.
- Application to different geological moments of Earth history (computing different effects of life)
- Application to any typology of exoplanet, connecting a disequilibrium characterization of a modeled planet/moon with its synthetic spectra. This will make the entropic analysis able to be compared with observed spectra.

Perspectives



Summary and outlook



- Calculate the extent of disequilibrium: no need of steady-state assumption
- KROME applied to disequilibrium calculations
- Reduce the selection of habitable planets to those which present high chemical disequilibrium (not attributable to any other process but life)
- Capability to separate the effect of each source of disequilibrium
- Life processes and photochemistry interact
- Mars disequilibrium is more than 10 orders of magnitude lower compared to present and pre-photosynthesis Earth
- Enhance precision for photochemical and early Earth's models

Thanks for your attention!



T. Grassi, Università di Roma - La Sapienza, Rome, Italy.

J. R. Brucato, Astrophysical Observatory of Arcetri - INAF, Firenze, Italy

S. O. Danielache, Sophia University, Tokio, Japan.

M. J. Russell, JPL, CalTech-NASA, Pasadena, CA, USA

S. Branciamore, Beckman Research Institute of City of Hope , CA, USA

A. Delgado-Bonal, J. Plá, T. Mendaza-de Cal, Maria Serrano, Centro de Astrobiologia, INTA-CSIC, Madrid, Spain

Kleidon, N. Virgo, Max Planck Institute for Biogeochemistry, Jena, Germany.

L. Grenfell, S. Gebauer, T.U. - DLR, Berlin, Germany.

.. and all members of the TDE Focus Group - NASA Astrobiology Institute